

Circulation of Eocene to Pliocene intermediate water masses in the Indian Ocean: Evidence from fossil fish teeth Nd isotopes

SCHER, H.D.¹, MARTIN, E.E.², HAASE, A.A.³

¹ Univ. of Florida, Gainesville, FL, 32611 (hscher@ufl.edu)

² Univ. of Florida, Gainesville, FL, 32611
(emartin@geology.ufl.edu)

³ Aires Consulting, Batavia, IL, 60510
(alisahaase@yahoo.com)

We have evaluated Nd and Sr isotopic compositions of fossil fish teeth from ODP site 757 (17°S, 88°E, 1650 m) on the Ninetyeast Ridge in the Indian Ocean. The calcareous nannofossil stratigraphy was enhanced with Sr age estimates from Oligocene forams. ϵ_{Nd} values of teeth increase from ~ -7.5 at 36 Ma to ~ -6 at 28.5 Ma. After 28.5 Ma values begin to decrease, falling to ~ -7.5 at 11.5 Ma before rising to a mid-Pleistocene value of -5.8 .

Variations in ϵ_{Nd} are believed to reflect changes in circulation of intermediate water masses in the Indian Ocean. The structure of the curve does not provide evidence of a Himalayan weathering input which is consistent with data from deeper sites in the Indian Ocean [3]. Today, site 757 is influenced by the Indonesian throughflow jet ($\epsilon_{Nd} \sim -5$), which injects Pacific waters into intermediate depths of the Indian Ocean. The low salinity core of AAIW ($\epsilon_{Nd} \sim -8$) is currently well developed south of 18°S in the Indian Ocean [1].

During the middle Eocene, $\delta^{18}O$ values at site 757 were higher than in the southern Indian Ocean, possibly reflecting a warm water mass overlying the [2]. The Tethys sea was a potential source for such a water mass (modern Mediterranean $\epsilon_{Nd} \sim -9$). Plate reconstructions for the Eocene show progressive constriction of the eastern Tethys connection with the Indian ocean. The shift toward radiogenic ϵ_{Nd} values from 35-28.5 Ma may reflect a decreased contribution of nonradiogenic Tethyan seawater, which was replaced by Indonesian throughflow water.

Crustal subsidence of site 757 into an AAIW-type water mass is one possible explanation for decreasing values from 28.5-11.5 Ma. Additionally, data from site 1090 (43°S, 20°W) suggest that waters sourced from the Southern Ocean were influenced by a nonradiogenic water mass during this interval, related to changes in circulation associated with the opening of the Drake Passage [4].

Values increase again after 11.5 Ma to -5.8 near the top of the core, approaching modern throughflow values.

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Lu-Hf ages of high pressure metamorphism in the Variscan fold belt of southern Germany

E. E. SCHERER, K. MEZGER, AND C. MÜNKER

Zentrallabor für Geochronologie, Institut für Mineralogie,
Universität Münster, Corrensstr. 24, D-48149 Germany
(escherer@nwz.uni-muenster.de)

The closure temperature of Lu-Hf in a particular garnet is ca. 50-150°C higher than that of Sm-Nd ([1], using the revised ^{176}Lu decay constant from [2]). Lu-Hf geochronology, in conjunction with Sm-Nd ages, is therefore a valuable tool for dating eclogite formation (i.e., garnet growth), or estimating exhumation rates. We have dated four Münchberg Massif (MM) eclogites with Lu-Hf on Grt-Omph pairs: two light eclogites from Oberkotzau (north-east MM), and two dark eclogites from Weissenstein (south-west MM). The dark variety (Cpx + Grt + Qtz + Rt \pm Hbl \pm Pheng) has a MORB-like bulk composition, whereas the light variety (Cpx + Grt + Qtz + Rt + Ky \pm Pheng \pm Zo \pm Hbl) has a high-Al basaltic composition [3, 4]. The light eclogites gave ages of 405 ± 7 Ma and 384 ± 6 Ma (2₋), and both have $_{-Hf}(t)$ values of +6.5. The dark eclogites yielded ages similar to those of the light variety, 397 ± 7 Ma and 398 ± 6 Ma, but have distinctively higher $_{-Hf}(t)$ of +12.2 and +11.9. Lu-Hf ages for both types overlap with published Sm-Nd and Rb-Sr internal isochron ages (380-395 Ma, [4]) for dark eclogites. The close agreement among all three isotope systems suggests extremely rapid exhumation soon after eclogitization. The high initial $_{-Nd}$ (ca. +8, [4]) and $_{-Hf}$ of the dark eclogites confirms that their mantle source was characterized by long-term depletion (e.g., MORB). In contrast, the lower initial $_{-Hf}$ and high Al-contents of the light eclogites may point to a subduction-related origin for their basaltic protoliths.

A suite of garnet metabasites from the Böllsteiner Odenwald (Mid-German Crystalline Rise) were once eclogites before being retrogressed to garnet amphibolites [5]. Our first attempt at dating the high pressure stage yielded Lu-Hf Grt-WR ages of 357 ± 7 Ma and 353 ± 11 Ma. Grt and wr from both samples plot on a single isochron (357 ± 6 Ma; MSWD = 1.09) and give an initial $_{-Hf}$ of +11.3, which is indicative of long term depletion in their mantle source. Some resetting may have occurred during retrograde metamorphism; the Lu-Hf ages are thus minimum ages for the eclogite facies event in the Odenwald.

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